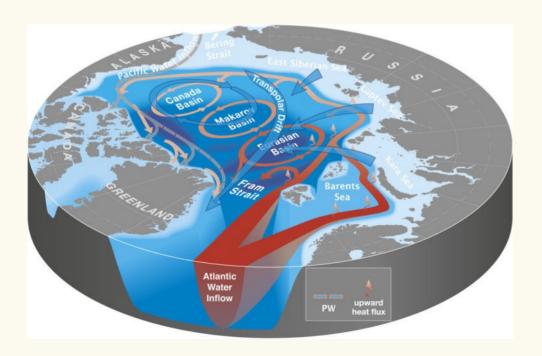
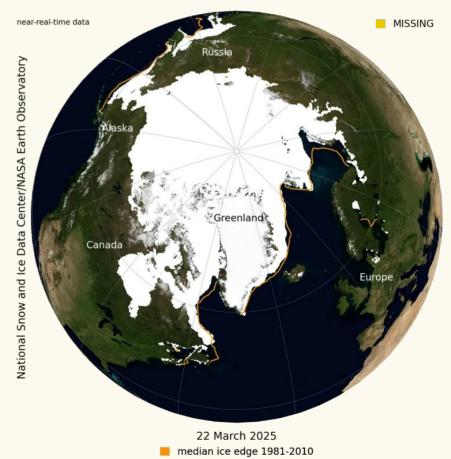
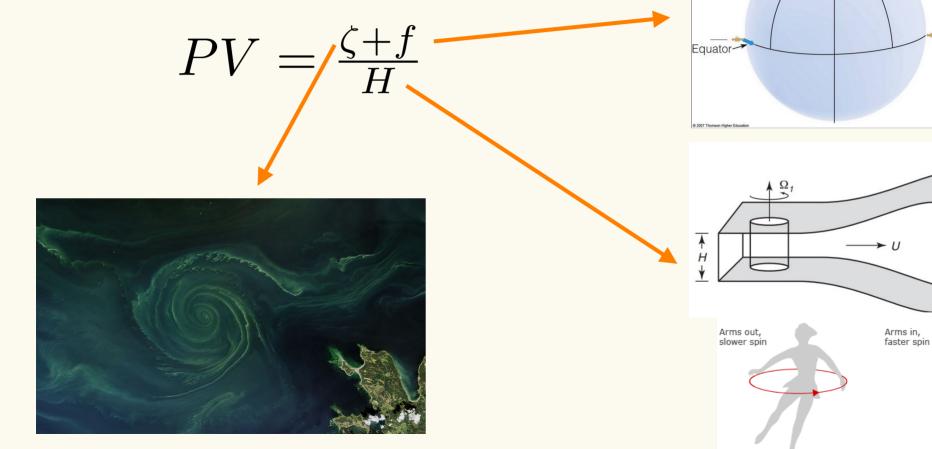
Idealized models for understanding Arctic Ocean Circulation

Anna Lina P. Sjur, Pål Erik Isachsen, Johan Nilsson, Joe LaCasce, Shaun Johnston, Magnus Dyrmose Ryseth





Potential vorticity



Maximum spin

Slight spin

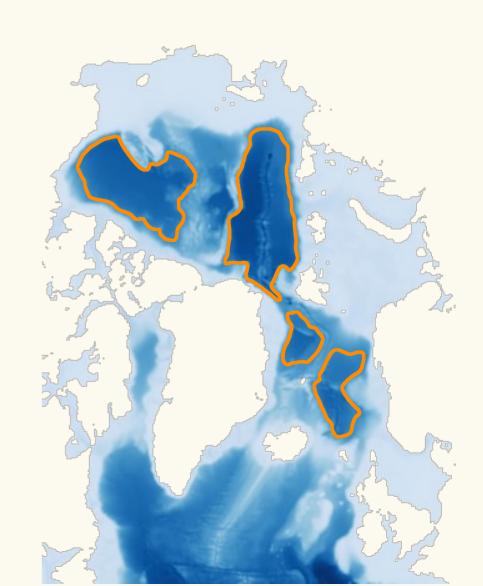
No spin

 $\Omega_2 > \Omega_1$

Potential vorticity

 Large scale currents have a strong tendency to follow contours of constant ambient potential vorticity.

These are the "railway tracks" of the ocean.

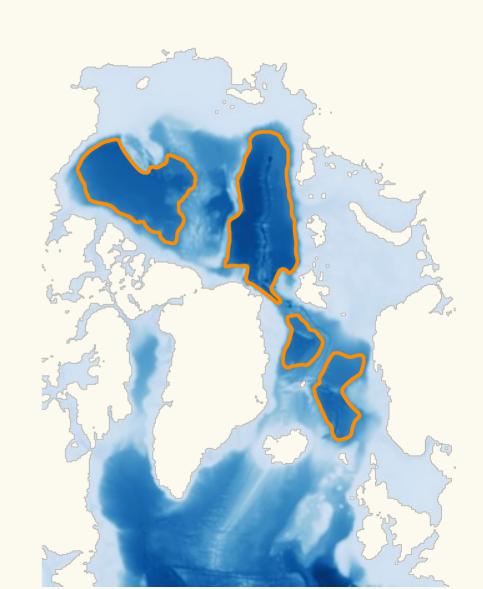


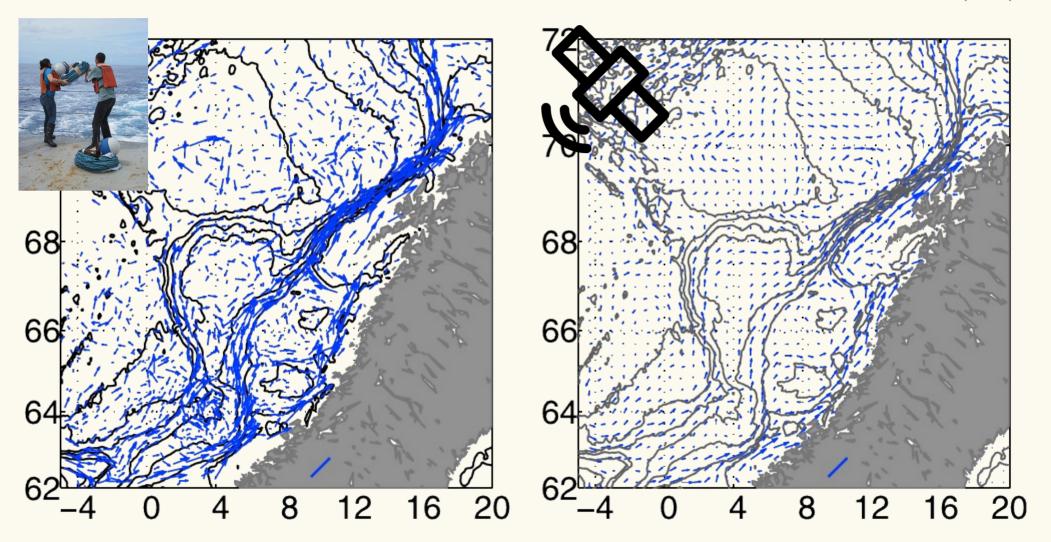
Potential vorticity

 Large scale currents have a strong tendency to follow contours of constant ambient potential vorticity.

These are the "railway tracks" of the ocean.

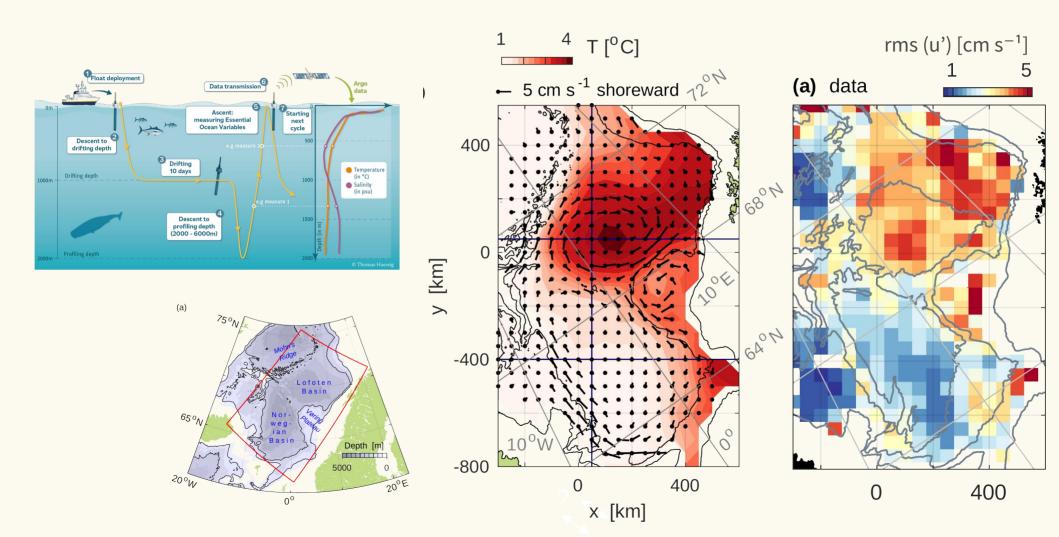
 Potential vorticity dominated by layer thickness at high latitudes → flow around closed depth contours.



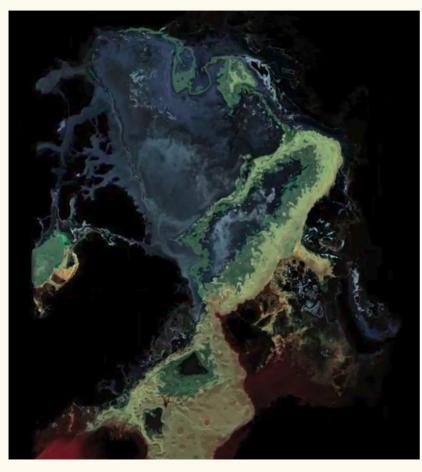


Observations from the Nordic Seas

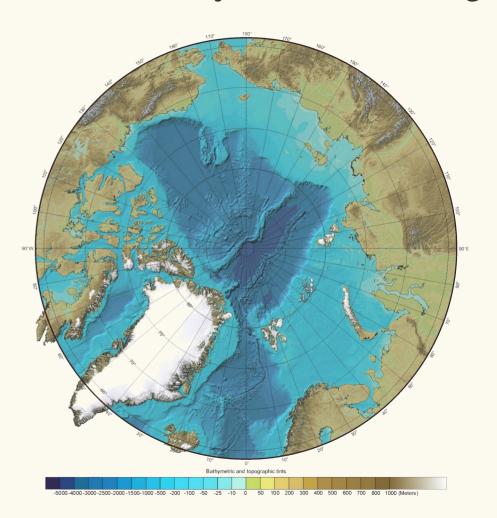
Johnston et al. (accepted)

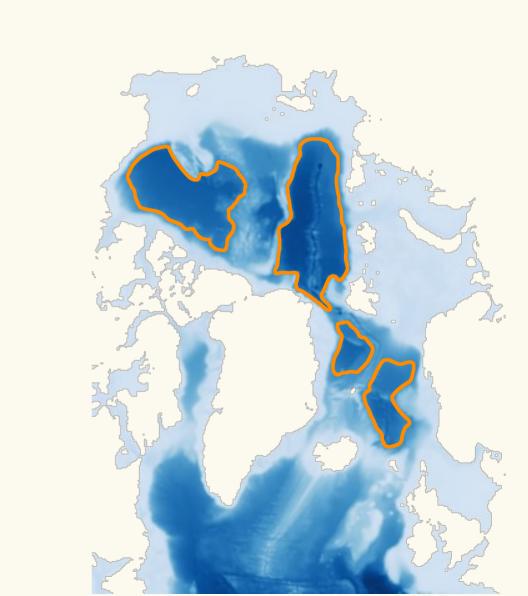


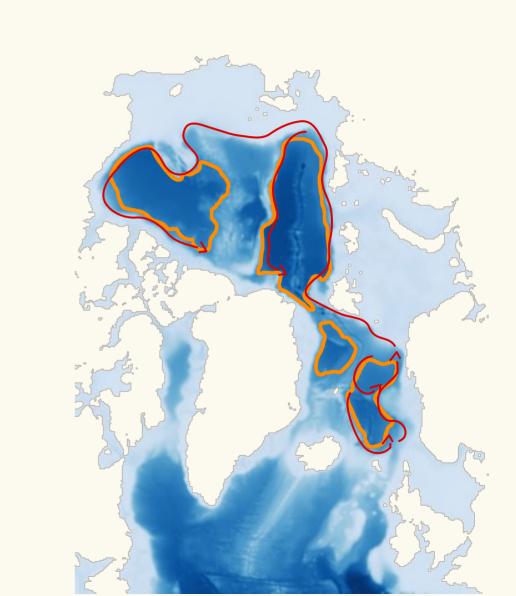
What about in the high Arctic? Rely on modeling

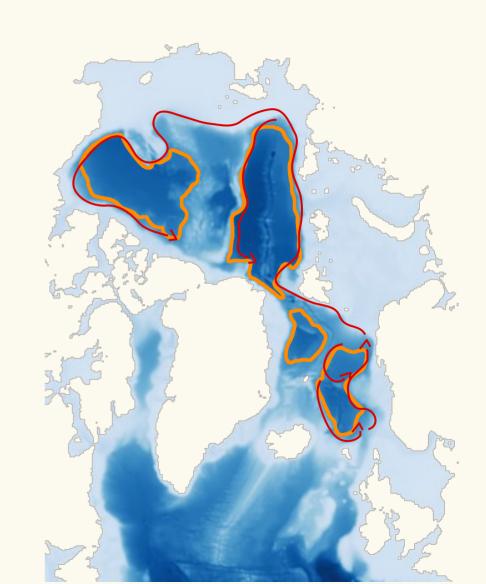


MITgcm, ~3.5 km resolution, TACC

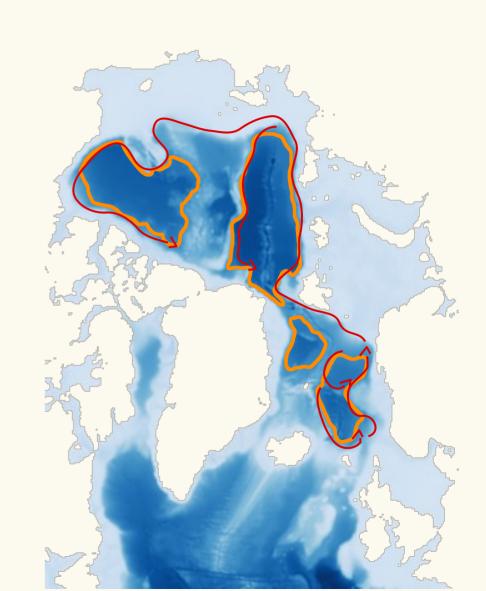








What can linear estimates teach us about the role of non-linearities?





Arctic4 CICE+ROMS





Linear model:

$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface Bottom circulation stress drag



JGR Oceans

RESEARCH ARTICLE

10.1029/2024JC021713

Key Points:

- · Time-varying ocean circulation around closed depth contours in simulations of the Arctic reflects low-pass filtered
- · A linear model, forced by surface stress and regulated by bottom friction, is able to reproduce the time-varying
- · The linear model lacks a cyclonic tendency, and has a slight asymmetry in response to the sign of surface

Correspondence to:

A. L. P. Sjur. a.l.p.sjur@geo.uio.no

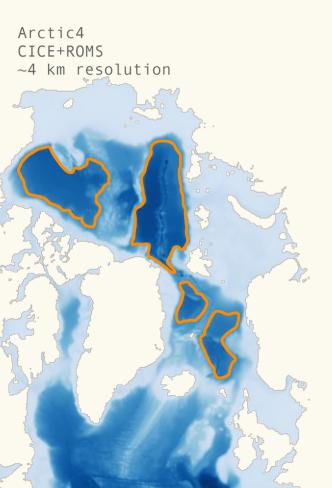
The Wind-Driven Time-Variable Circulation in the Arctic Mediterranean

A. L. P. Sjur¹, P. E. Isachsen^{1,2}, J. Nilsson³, J. H. LaCasce¹, and M. D. Ryseth^{1,4}

¹University of Oslo, Oslo, Norway, ²Norwegian Meteorological Institute, Oslo, Norway, ³Stockholm University, Stockholm, Sweden, 4Statkraft Energi AS, Oslo, Norway

Abstract The Arctic Ocean is a key component of Earth's climate system, and an understanding of ocean dynamics in this region is central for predicting how the Arctic is responding to a changing climate. In this study, we examine the ocean circulation in a high-resolution numerical model of the Arctic Ocean and Nordic Seas. Based on what is observed in this simulation, we reexamine an existing idealized linear model estimating the time-variable large-scale circulation in ocean basins, and test it against the highly nonlinear numerical model. The idealized model is an integral relation derived from the linear momentum equations and assumes that the circulation around a closed depth contour is driven by surface stresses and regulated by bottom friction. We show that the idealized model estimates agree very well with the numerical simulations. This indicates that much of the variability of the large-scale circulation can be explained by linear processes. In particular, a correct description of the net surface stress over partially ice-covered areas improves the correlation between linear model and numerical simulations significantly in the Arctic Ocean compared to a previous study. However





Linear model:

$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface Bottom circulation stress drag

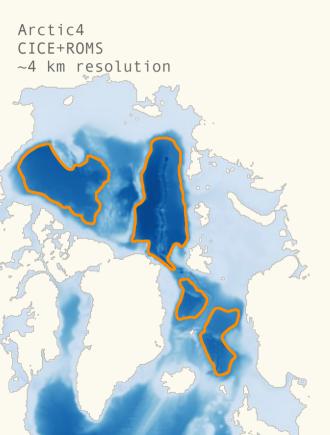
 $\oint \boldsymbol{u} \cdot d\boldsymbol{l} = \exp\left(-\frac{R}{H}t\right) \oint \boldsymbol{u}^{0} \cdot d\boldsymbol{l} + \int_{t_{0}}^{t} \exp\left[-\frac{R}{H}\left(t - t'\right)\right] \frac{1}{H} \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} dt'$

Circulation

Decaying contribution from initial state

Filtered surface stress



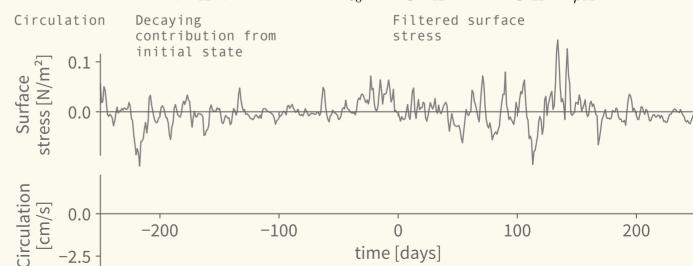


Linear model:

$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface Bottom circulation stress drag

$$\oint \boldsymbol{u} \cdot d\boldsymbol{l} = \exp\left(-\frac{R}{H}t\right) \oint \boldsymbol{u}^{0} \cdot d\boldsymbol{l} + \int_{t_{0}}^{t} \exp\left[-\frac{R}{H}\left(t - t'\right)\right] \frac{1}{H} \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} dt'$$





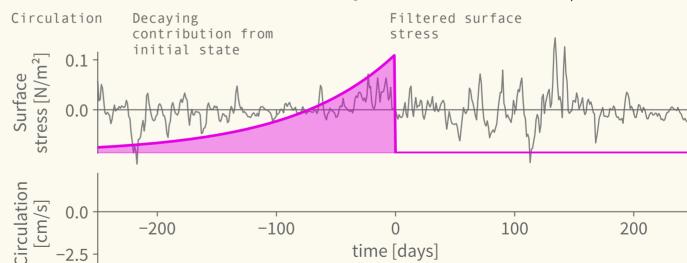


Linear model:

$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface Bottom circulation stress drag

$$\oint \boldsymbol{u} \cdot d\boldsymbol{l} = \exp\left(-\frac{R}{H}t\right) \oint \boldsymbol{u}^0 \cdot d\boldsymbol{l} + \int_{t_0}^t \exp\left[-\frac{R}{H}\left(t - t'\right)\right] \frac{1}{H} \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} dt'$$





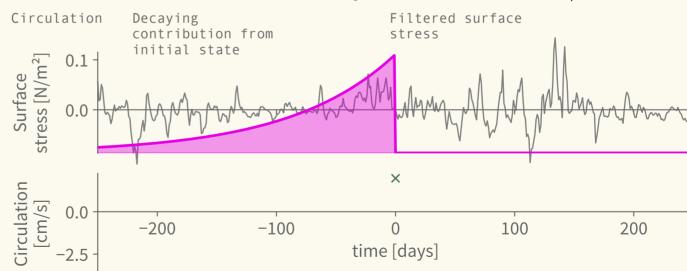


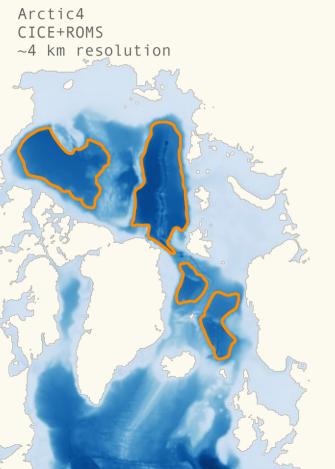


$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface Bottom circulation stress drag

$$\oint \boldsymbol{u} \cdot d\boldsymbol{l} = \exp\left(-\frac{R}{H}t\right) \oint \boldsymbol{u}^{0} \cdot d\boldsymbol{l} + \int_{t_{0}}^{t} \exp\left[-\frac{R}{H}\left(t - t'\right)\right] \frac{1}{H} \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} dt'$$









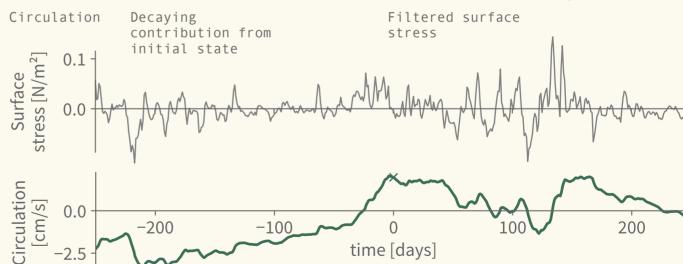
Linear model:

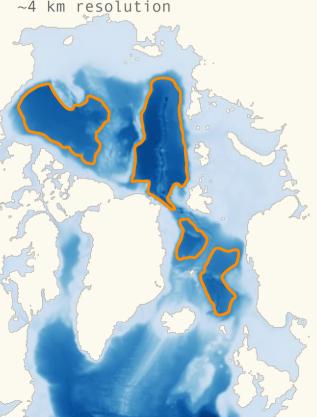
$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l}$$

Change in Surface circulation stress

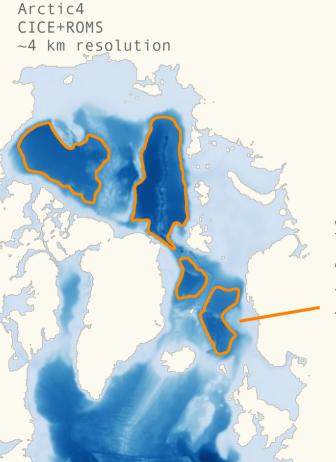
Bottom drag

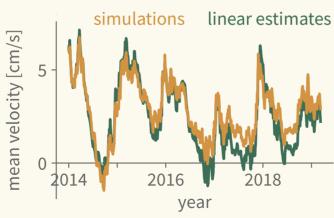
$$\oint \boldsymbol{u} \cdot d\boldsymbol{l} = \exp\left(-\frac{R}{H}t\right) \oint \boldsymbol{u}^{0} \cdot d\boldsymbol{l} + \int_{t_{0}}^{t} \exp\left[-\frac{R}{H}\left(t - t'\right)\right] \frac{1}{H} \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} dt'$$

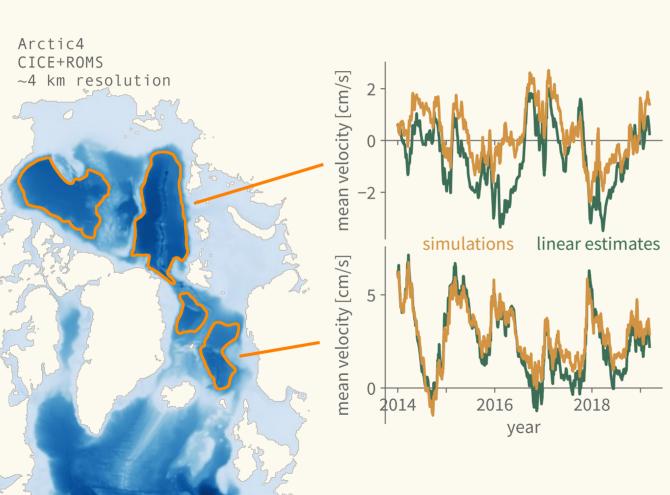




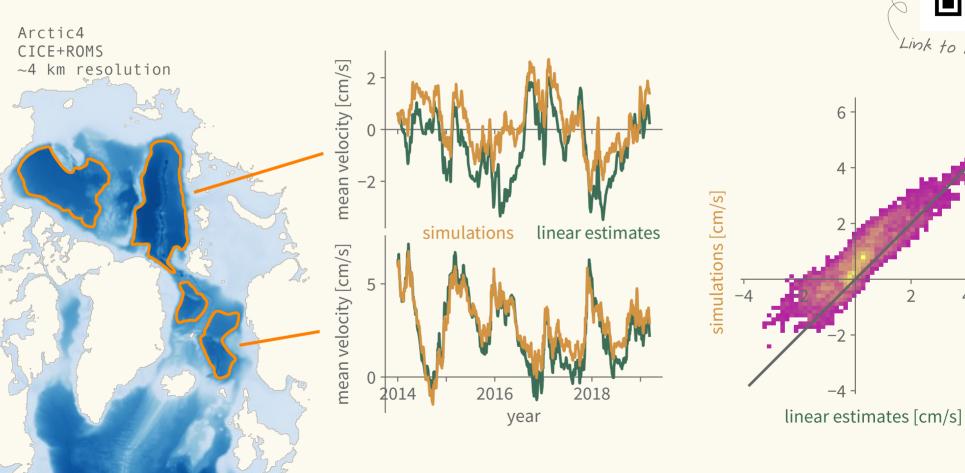




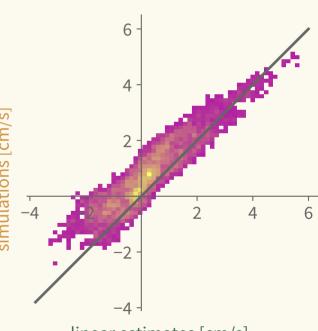


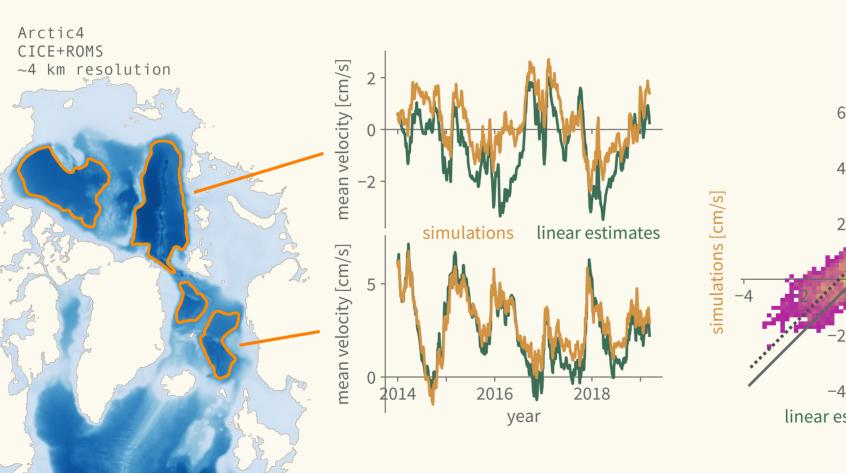


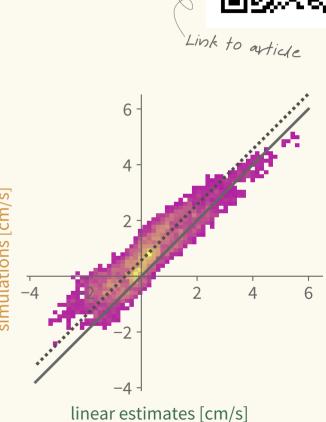




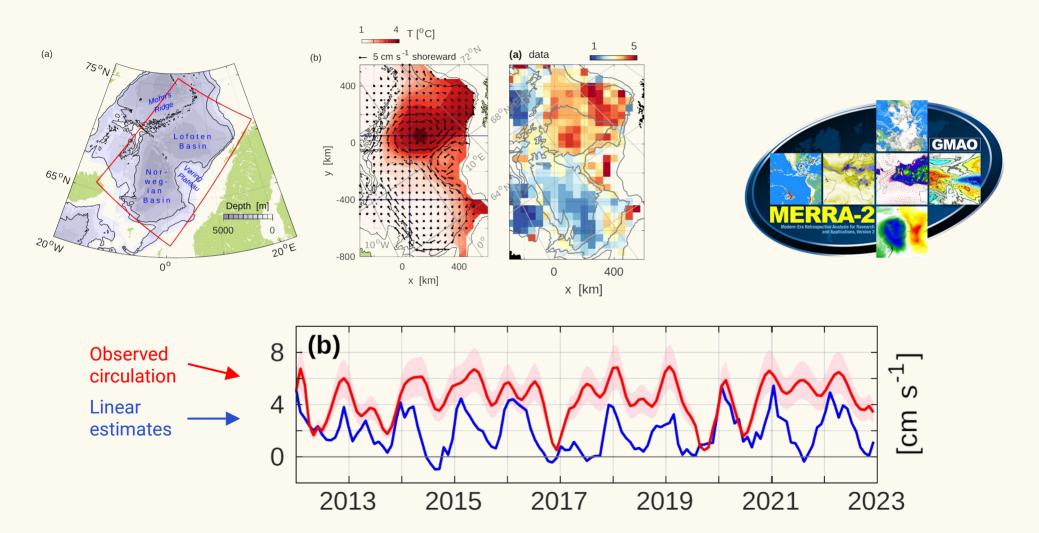








Comparison with Lofoten Basin observations (accepted)



What enhances cyclonic circulation?

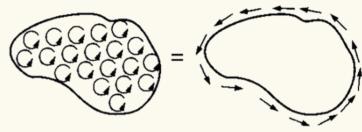
On depth contours, the main suspect is flux of vorticity

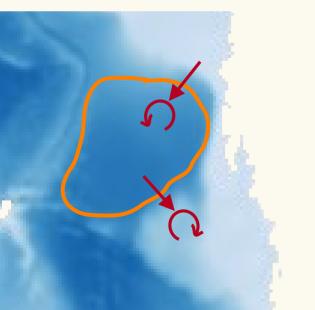
 $\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l} - \oint \overline{\zeta \boldsymbol{u}} \cdot \boldsymbol{n} \, dl$

Change in circulation

Surface stress

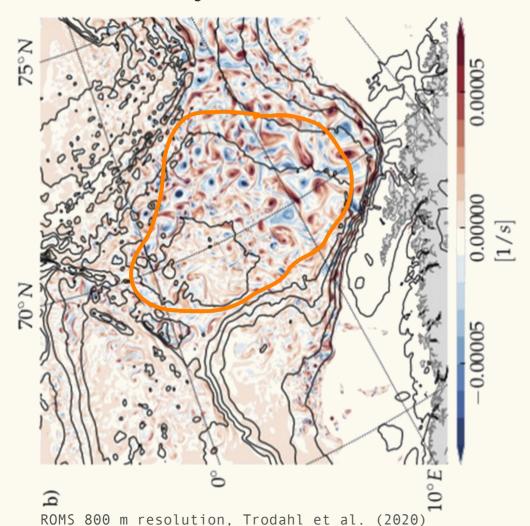
Bottom drag Flux of vorticity Green's theorem







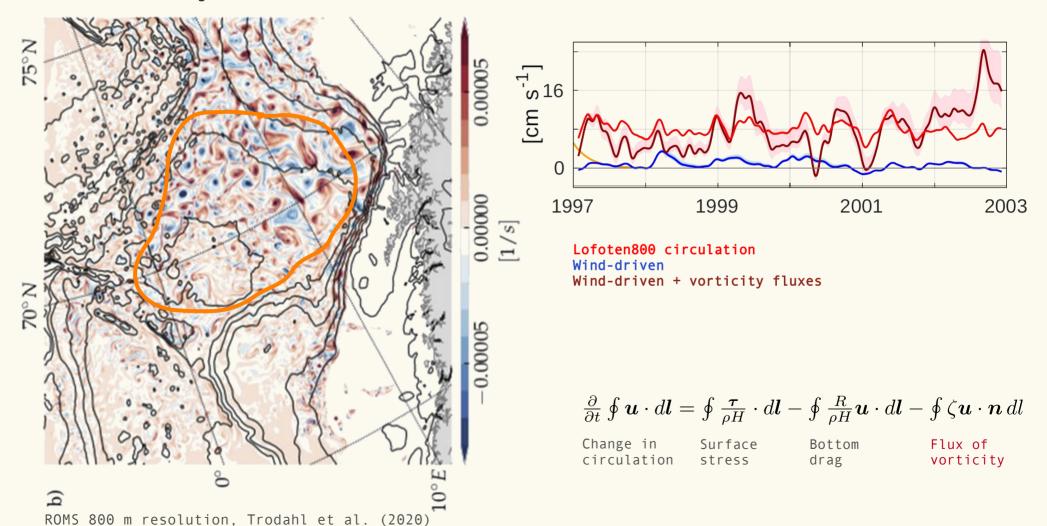
Vorticity flux contribution in simulations



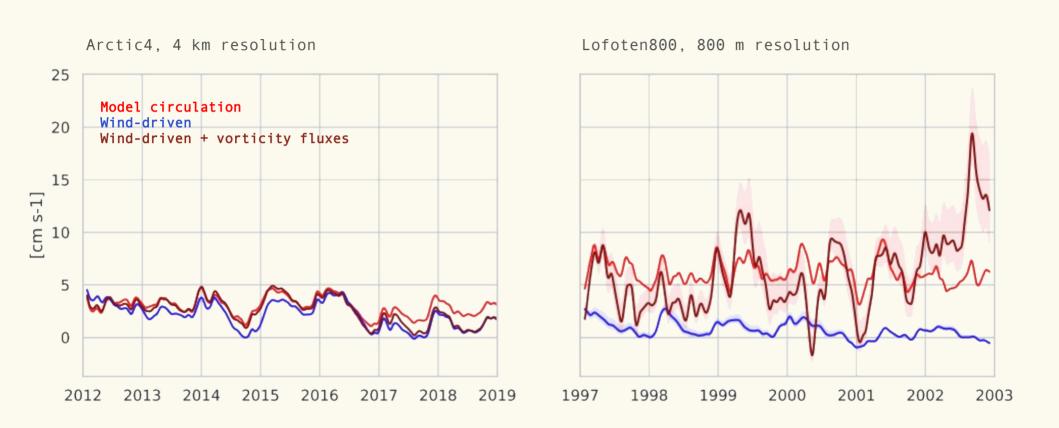
$$\frac{\partial}{\partial t} \oint \boldsymbol{u} \cdot d\boldsymbol{l} = \oint \frac{\boldsymbol{\tau}}{\rho H} \cdot d\boldsymbol{l} - \oint \frac{R}{\rho H} \boldsymbol{u} \cdot d\boldsymbol{l} - \oint \zeta \boldsymbol{u} \cdot \boldsymbol{n} \, dl$$

Change in Surface Bottom Flux of circulation stress drag vorticity

Vorticity flux contribution in simulations



The need for high resolution

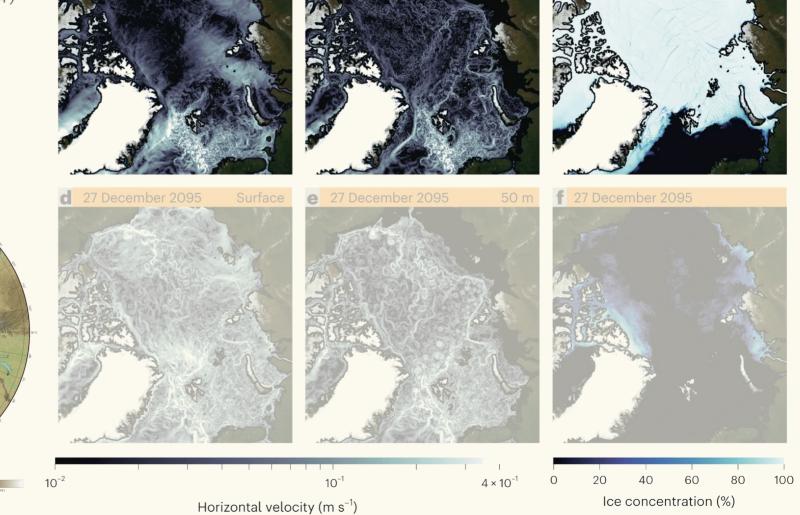


Blue Arctic

27 December 2015

Surface

Li et al. Nature C. (2024) FESOM2, 1 km resolution



b 27 December 2015

c 27 December 2015

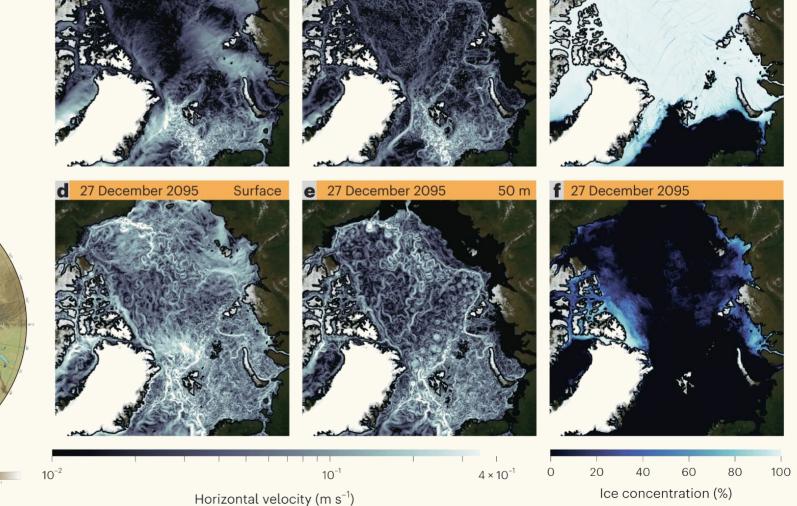
50 m

Blue Arctic

27 December 2015

Surface

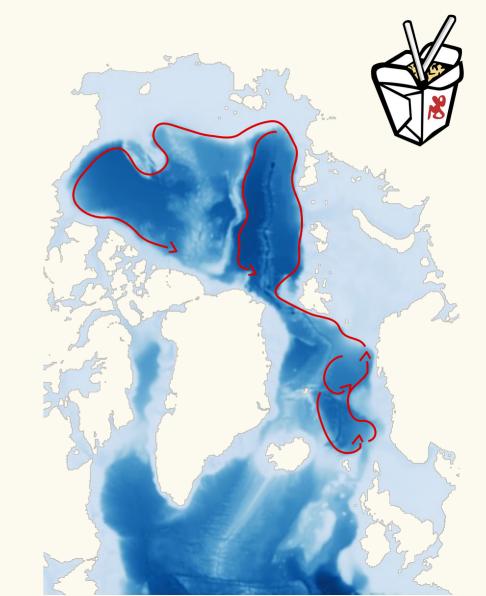
Li et al. Nature C. (2024) FESOM2, 1 km resolution



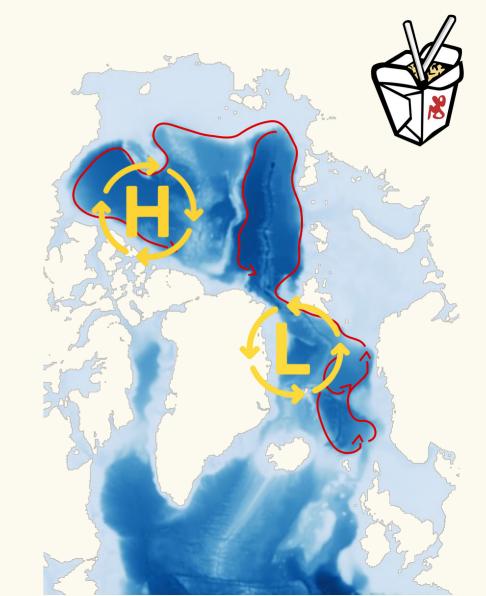
b 27 December 2015

c 27 December 2015

50 m

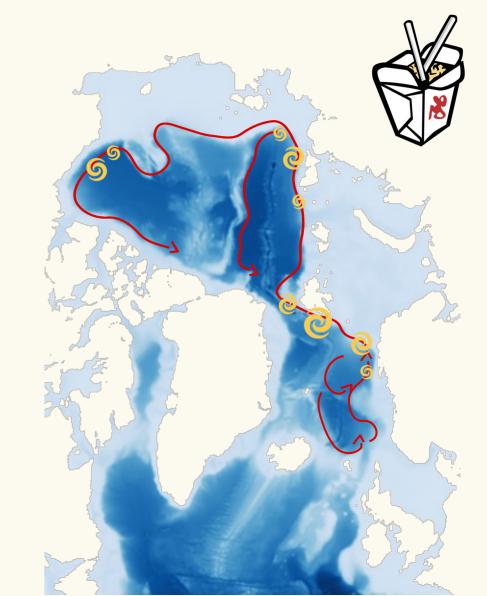


Circulation responds, to a large degree, linearly to surface stress.



Circulation responds, to a large degree, linearly to surface stress.

However, vorticity fluxes connected to eddy activity gives a positive shift.



Circulation responds, to a large degree, linearly to surface stress.

However, vorticity fluxes connected to eddy activity gives a positive shift.

Very high resolution is needed to properly simulate positive shift.

